

Warming alters coupled carbon and nutrient cycles in experimental streams

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Abstract

Although much effort has been devoted to quantifying how warming alters carbon cycling across diverse ecosystems, less is known about how these changes are linked to the cycling of bioavailable nitrogen and phosphorus. In freshwater ecosystems, benthic biofilms (i.e. thin films of algae, bacteria, fungi, and detrital matter) act as biogeochemical hotspots by controlling important fluxes of energy and material. Understanding how biofilms respond to warming is thus critical for predicting responses of coupled elemental cycles in freshwater systems. We developed biofilm communities in experimental streamside channels along a gradient of mean water temperatures (7.5–23.6 °C), while closely maintaining natural diel and seasonal temperature variation with a common water and propagule source. Both structural (i.e. biomass, stoichiometry, assemblage structure) and functional (i.e. metabolism, N₂-fixation, nutrient uptake) attributes of biofilms were measured on multiple dates to link changes in carbon flow explicitly to the dynamics of nitrogen and phosphorus. Temperature had strong positive effects on biofilm biomass (2.8- to 24-fold variation) and net ecosystem productivity (44- to 317-fold variation), despite extremely low concentrations of limiting dissolved nitrogen. Temperature had surprisingly minimal effects on biofilm stoichiometry: carbon:nitrogen (C:N) ratios were temperature-invariant, while carbon:phosphorus (C:P) ratios declined slightly with increasing temperature. Biofilm communities were dominated by cyanobacteria at all temperatures (>91% of total biovolume) and N₂-fixation rates increased up to 120-fold between the coldest and warmest treatments. Although ammonium-N uptake increased with temperature (2.8- to 6.8-fold variation), the much higher N₂-fixation rates supplied the majority of N to the ecosystem at higher temperatures. Our results demonstrate that temperature can alter how carbon is cycled and coupled to nitrogen and phosphorus. The uncoupling of C fixation from dissolved inorganic nitrogen supply produced large unexpected changes in biofilm development, elemental cycling, and likely downstream exports of nutrients and organic matter.

Keywords: benthic biofilms, climate change, coupled biogeochemical cycles, ecological stoichiometry, metabolism, N₂-fixation, nutrient uptake, temperature

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Introduction

Global mean air temperatures rose ~0.89 °C in the twentieth century, with demonstrated effects on ecological structure and function across Earth's ecosystem types (Parmesan & Yohe, 2003; IPCC, 2013). Like their terrestrial counterparts, freshwater ecosystems have shown clear warming trends (Adrian *et al.*, 2009;

Kaushal *et al.*, 2010), with associated changes in the physiology of individuals, structure of communities, and rates of ecosystem processes (Woodward *et al.*, 2010). Considerable effort has been devoted to understanding how such warming is likely to alter carbon (C) cycling (Cox *et al.*, 2000; Yvon-Durocher *et al.*, 2010a) because the balance of anabolic (i.e. production) and catabolic (i.e. respiration) processes can drive feedbacks that may either exacerbate or slow the accumulation of CO₂ in the atmosphere (Luo *et al.*, 2004; Yvon-Durocher & Allen, 2012). Much less attention, however, has been given to how C-cycle responses to warming may influence the flux and storage of other biologically important elements, such as nitrogen (N) and phosphorus (P), the cycling of which is coupled to C (Schlesinger *et al.*,

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2011). Recent models and empirical work demonstrate that these coupled elements cannot be ignored (Hungate *et al.*, 2003; Finzi *et al.*, 2011; Cross *et al.*, 2015).

How shifts in the C cycle influence the dynamics of coupled nutrients should depend on how the demand for these nutrients changes with warming through (i) species-level physiological adjustments and/or (ii) patterns of species dominance and altered rates of ecosystem processes (Sistla & Schimel, 2015). Predicted increases in C fixation rates with warming (e.g. Yvon-Durocher *et al.*, 2010a) should drive changes in N and P demand that broadly reflect the C:nutrient stoichiometry (and potential demand) of primary producers. Because the elemental composition of autotrophs is potentially plastic, warming of nutrient-limited environments may lead to elevated autotroph C:nutrient stoichiometry (Frost *et al.*, 2005; Sardans *et al.*, 2012; De Senerpont Domis *et al.*, 2014; *but see* Hall *et al.*, 2005; Nifong *et al.*, 2014), reduced relative nutrient demand, and increased nutrient use efficiency (UE; Rhee & Gotham, 1981; Cross *et al.*, 2015). Given no change in community structure, such stoichiometric responses to warming may modify how autotrophic C flux (e.g. gross primary production) is linked to the uptake and cycling of bioavailable N and P (e.g. nitrate- and ammonium-nitrogen and phosphate-phosphorus in aquatic systems). Such an effect is also predicted if warming leads to broad shifts in community structure that favor taxa with elevated C:nutrient ratios, reduced nutrient demand, and relatively high nutrient UE (Cross *et al.*, 2015).

Warming-induced changes in the activity and structure of autotrophic aquatic communities may also drive fundamental shifts in how nutrients are obtained from the environment and cycled. In N-limited ecosystems, warming can favor the abundance and activity of dinitrogen (N₂)-fixing cyanobacteria, which typically exhibit high optimum growth temperatures (Paerl & Huisman, 2008; Scott & Marcarelli, 2012; *but see* Lürling *et al.*, 2013) and steep temperature dependence for N₂-fixation (Ceuterick *et al.*, 1978). These taxa can obtain N from gaseous N₂, effectively decoupling C-based primary production from the dissolved ionic forms of N (e.g., ammonium, nitrate) utilized by non-N₂-fixing taxa. If such changes in the source and routing of nutrients are ubiquitous, our understanding of how warming influences coupled element cycles will need significant revision (Finzi *et al.*, 2011).

Here, we quantify the effects of warming on the coupling of C, N, and P cycling in experimental stream channels. Our small-scale experimental system provided a tractable framework for quantifying warming responses of stream biofilms, which represent hotspots of biogeochemical activity and contribute substantially to whole-system nutrient cycling

and energy flux in lotic ecosystems (Lock *et al.*, 1984; Battin *et al.*, 2003). We first examined how warming influenced the biomass and elemental composition of biofilms. We predicted that warming would increase C-fixation rates, but reduce total biomass, consistent with metabolic scaling patterns (Brown *et al.*, 2004; Allen *et al.*, 2005; Yvon-Durocher *et al.*, 2010b). We also predicted that warming this N-limited system (Friberg *et al.*, 2009) would increase C:nutrient ratios, mediated by physiological adjustments within the community and elevated nutrient UEs. Alternatively, if warming enhances N₂-fixation, eliminating progressive N limitation (Hungate *et al.*, 2003), we would predict little change in biofilm C:nutrient ratios. Such a result would be driven by N₂-fixation (and flux of N to the biofilm) keeping pace with C-fixation as temperatures increase.

We also examined how changes in biofilm biomass and stoichiometry were associated with relative rates of ecosystem processes (i.e. gross primary production, ecosystem respiration, N and P uptake, and N₂ fixation) and element coupling. We predicted that temperature dependent patterns of N and P uptake would reflect shifts in biofilm stoichiometry, with little change across the thermal gradient as a result of increased C:nutrient ratios. However, if N₂-fixers are favored by warming, we expected that N demand by autotrophs would shift from heavy reliance on dissolved inorganic forms of N to dissolved N₂ gas, effectively decoupling C-fixation from bioavailable inorganic N. This could translate to previously unpredicted C cycling dynamics under future, presumably warmer, climates.

Materials and methods

Site description

Our study was conducted in the geothermally active Hengill area of southwestern Iceland (64.056°N, 21.283°W). This region contains a network of small streams and springs that vary significantly in ambient temperature (~6–100 °C). These large temperature differences result from indirect, local geothermal heating of soils and bedrock, and streams consequently show little variation in solute chemistry, which is similar to many pristine subarctic streams. Importantly, indirect geothermal heating does not produce unusual solute chemistry often associated with volcanic regions (Friberg *et al.*, 2009; O'Gorman *et al.*, 2012). The dominant geology is young basalt (Árnason *et al.*, 1969). Streams in the study area exhibit low dissolved molar N:P ratios (<3), with primary production believed to be primarily, but not exclusively, N-limited (Friberg *et al.*, 2009; Gundmundsdottir *et al.*, 2011).

Experimental design

We conducted a controlled and replicated warming experiment using a relatively cold (6.7 ± 1.7 °C, mean \pm SD during May–Aug 2013) unnamed tributary of the Hengladalsá River. We manipulated water temperature in our treatments (Table 1) using gravity-fed heat-exchanger systems (HEX) placed in geothermally heated springs (see Welter *et al.*, 2015 for full design description). The HEX supplied water of varying temperatures (7.5–23.6 °C) to 15 open-topped U-shaped PVC channels (length: 3 m, width: 25.4 mm, height: 40 mm) arrayed in parallel facing magnetic south (Fig. S2). Our temperature treatments were chosen to complement the temperature range (~6–30 °C) of natural streams in the Hengill region (i.e. streams without extremophilic organisms). Channels were divided into three blocks with five channels per block. Temperature treatments were randomly assigned to channels within blocks ($n = 5$ treatments, $n = 3$ replicates). Water temperature at the head of each channel was logged (HOBO Pendant, Onset Computer Corp., MA, USA) at 15-min intervals. Mean channel discharge, velocity, and depth were standardized at 0.03 ± 0.004 L/s, 0.24 ± 0.05 m/s, and 6.1 ± 1.7 mm (mean \pm 1 SD) across channels respectively. Preleached basalt tiles (25 \times 25 mm; Deko Tile, CA, USA; $n \approx 110$ per channel) were deployed on 20 May 2013, and biofilm communities were allowed to colonize tiles naturally. All tiles were destructively sampled without replacement throughout the experiment.

Ecosystem metabolism

We assessed ecosystem metabolism on tiles three times (after 30, 42, and 58 days) by measuring oxygen changes in 0.27 L clear, recirculating chambers (Welter *et al.*, 2015). On each date, four tiles were randomly selected from each treatment (total surface area 24 cm²) and placed in chambers containing sieved (250- μ m) HEX source water from each treatment. We measured net ecosystem production (NEP) under ambient light conditions and ecosystem respiration (ER) in complete darkness, maintaining consistent treatment temperatures. Dissolved oxygen (DO) and internal chamber temperatures were recorded at 1-min intervals (YSI, Pro-ODO, OH, USA) during the ~1.5 h incubations. NEP and ER ($\text{mg DO m}^{-2} \text{ h}^{-1}$) were calculated as:

$$(\text{NEP}, \text{ER}) = (\Delta\text{O}_2 - \text{WC}) \times (V/S)$$

where ΔO_2 is the change in DO concentration per unit time ($\text{mg DO L}^{-1} \text{ h}^{-1}$), V is chamber volume (L), S is the active surface

area of the tiles (m²), and WC is the water column metabolic rate ($\text{mg DO L}^{-1} \text{ h}^{-1}$), which was minimal. Gross primary production (GPP) was calculated as $\text{GPP} = \text{NEP} + \text{ER}$ (Bott, 2006).

Nutrient uptake

We measured $\text{NH}_4\text{-N}$ and SRP uptake in the recirculating chambers on three dates ($\text{NH}_4\text{-N}$: 31, 44, and 65 days; SRP: 32, 43, and 66 days) on different tiles (total surface area 24 cm²). Ammonium-N and orthophosphate-P spikes (as NH_4Cl and Na_2HPO_4) were added to a target concentration of ~50 $\mu\text{g/L}$ above ambient (~5-fold N and ~3-fold P increase) at the beginning of the incubations. This level of nutrient addition was chosen to reduce the risk of nutrient-limiting conditions during incubations. We believe this approach captured values that are largely representative of natural uptake rates; however, an addition of surplus bioavailable nutrients may have resulted in slightly overestimated uptake rates (O'Brien & Dodds, 2008). Triplicate water samples were collected at the beginning and end of incubations (typically 1.5 h), filtered immediately through 0.45- μm glass-fiber filters and stored <5 h in the dark until frozen. We analyzed samples with colorimetric (SRP; APHA, 2012) or fluorometric techniques ($\text{NH}_4\text{-N}$; Holmes *et al.*, 1999; Taylor *et al.*, 2007). Areal uptake ($\mu\text{g m}^{-2} \text{ h}^{-1}$) was calculated as the difference between nutrient mass ($\mu\text{g L}^{-1} \times$ chamber volume [L]) pre- and postincubation, divided by the surface area (m²) of tiles and incubation time (h).

N_2 -fixation rates

N_2 -fixation was measured on three occasions (41, 53, and 69 days) with acetylene reduction assays in the recirculating chambers (Flett *et al.*, 1976; Capone, 1993). Incubation length was ~2 h, with tile selection and handling as outlined above. Gas samples were analyzed with a gas chromatograph (SRI Instruments, 8610, Torrance, CA, USA) within 72 h of collection, and areal uptake rates of N_2 -fixation were calculated as the difference between pre- and postincubation gas concentrations divided by tile surface area (m²) and length of incubation (h), assuming a 3:1 acetylene: N_2 conversion ratio (Capone, 1993; Welter *et al.*, 2015).

Biofilm structure and stoichiometry

We scrubbed biofilms from tiles immediately following metabolism and N_2 -fixation incubations, and the resultant slurry

Table 1 Physicochemical characteristics of each temperature treatment

Treatment	Mean temperature	$\text{NH}_4\text{-N}$	$\text{NO}_3\text{-N}$	SRP	DO	Specific conductivity
A	7.5 (2.3)	11.6 (7.1) ^a	52.2 (44.2) ^a	16.6 (7.8) ^a	11.5 (0.4) ^a	68.8 (2.6) ^a
B	11.2 (2.2)	7.0 (4.7) ^a	37.0 (31.3) ^a	19.0 (6.8) ^a	11.5 (0.4) ^a	67.7 (2.8) ^{ab}
C	15.5 (2.3)	11.5 (5.5) ^a	46.0 (40.6) ^a	15.3 (4.0) ^a	11.4 (0.5) ^a	67.4 (2.7) ^{ab}
D	19.0 (2.2)	7.5 (5.1) ^a	45.2 (44.3) ^a	18.6 (8.7) ^a	10.7 (0.5) ^b	67.4 (3.0) ^c
E	23.6 (2.4)	11.8 (7.3) ^a	41.2 (43.2) ^a	15.4 (7.3) ^a	10.0 (0.4) ^b	67.6 (2.8) ^d

Mean treatment temperature (°C \pm 1 SD), ammonium-N ($\mu\text{g N/L} \pm 1$ SD), nitrate ($\mu\text{g N/L} \pm 1$ SD), soluble reactive phosphorus (SRP; $\mu\text{g P/L} \pm 1$ SD), dissolved oxygen (DO; $\text{mg O}_2/\text{L} \pm 1$ SD), and specific conductivity ($\mu\text{S} \pm 1$ SD). Superscript letters indicate statistically similar groups.

was homogenized, subsampled and filtered onto glass-fiber filters to estimate ash-free dry mass (AFDM), C, N, and P content (Whatman™ GF/F filters), and chlorophyll *a* (chl *a*; Andwin Scientific A/E filters). Individual and separate filters were used for all response metrics, except for C and N, which were processed on the same filter. We only sampled nutrient incubation tiles for AFDM. Chlorophyll *a* samples were collected in amber bottles and kept cool during transport to our laboratory, where they were processed immediately. We used a ~24 h acetone extraction and measured chl *a* with a handheld fluorometer (Turner Designs, Aquafluor, CA, USA). AFDM, C, N, and P filters were dried at 55 °C for ≥72 h. AFDM reflects the difference between filter dry weight and ash weight after 2 h at 500 °C. We determined C and N content with an elemental analyzer (Costech Analytical Technologies, Inc., CA, USA) and P content with colorimetric techniques (American Public Health Association (APHA), 2012). Community composition subsamples were preserved in ~5% glutaraldehyde, concentrated via centrifugation, and examined at 400 × magnification in a 0.1-mL Palmer counting chamber (Lowe & LaLiberte, 2006). Biovolume of distinct autotroph taxa (largely identified to genus) was calculated using established taxon-specific geometries (Sun & Liu, 2003).

Nutrient demand and use efficiency

We calculated a predicted net nutrient demand of biofilms to assess whether our estimates of predicted demand matched our real measurements of N and P uptake rate. Predicted net nutrient demand was calculated as the product of measured NEP (moles C m⁻² h⁻¹) and biofilm nutrient:C ratio. NEP values were converted to units of C assuming a photosynthetic quotient of 1.2 (Bott, 2006). We chose NEP over the more commonly used NPP (Pastor & Bridgman, 1999) to avoid assumptions about the fraction of GPP respired by autotrophs.

To assess how temperature influenced the coupling of C and nutrient fluxes, we calculated biofilm nutrient UEs for N and P across the thermal gradient. We present two measures of nutrient UE. First, we present biofilm C:N and C:P ratios, which can reflect the net nutrient UE of the community if we assume that (i) biofilms are dominated by actively growing cells and (ii) biofilm stoichiometry reflects net nutrient requirements over ecologically important time scales. Second, we calculated biofilm nutrient UEs using our measured fluxes of C and nutrients into the biofilm. Here, nutrient UE was calculated as mean NEP (moles C m⁻² h⁻¹) per unit nutrient acquired (moles N or P m⁻² h⁻¹; Pastor & Bridgman, 1999). P use efficiency (P-UE) was calculated with SRP uptake rate, while nitrogen use efficiency (N-UE) was calculated using NH₄⁺-N uptake rates only (N-UE_I) or with combined N₂-fixation and NH₄⁺-N uptake rates (N-UE_{UF}).

Data analysis

We analyzed data with repeated measures mixed-effects models fit with the 'lme' function in the R package 'nlme'. Correlation coefficients were generated with the 'r.squaredGLMM' function in the 'MUMIN' package (R Core Team, 2013). We fit an initial 'full' model that included temperature, day, and a

temperature × day interaction as fixed effects, and channel identity as a random effect. We used model selection techniques (Akaike Information Criterion; AIC; Bozdogan, 1987) to determine the most appropriate model given the data. Occasionally, residual plots suggested nonlinear responses to temperature. For these metrics we used repeated measures analysis of variance (RMANOVA) and AIC value model selection (Bozdogan, 1987; Cottingham *et al.*, 2005). This is a more conservative approach for addressing potential threshold temperature effects, thus avoiding assumptions about the shape of nonlinear models. Least-squares regressions were fit to individual sample periods with the 'lm' function in the R package 'stats' when limited data points restricted use of mixed-effects models (R Core Team 2013). Individual rate measurements were discarded when below our detection thresholds. Response metrics were log_e(*x*)-transformed to meet model assumptions of normality when appropriate, and all analyses were conducted with an alpha of 0.05.

Results

Temperature manipulation

The ambient temperature treatment had a temperature regime that was largely similar – in both temperature and diel pattern – to the source stream. The other four treatments were on average 3.7, 8.0, 11.5, and 16.1 °C above ambient, while also exhibiting diel temperature patterns similar to the source stream (Table 1; Fig. S1). Dissolved nutrient concentrations at the upstream end of the channels did not differ among treatments (RMANOVA: SRP, $F_{5,20} = 0.85$, $P = 0.53$; NH₄⁺-N, $F_{5,20} = 2.06$, $P = 0.113$; NO₃-N, $F_{5,20} = 0.67$, $P = 0.65$). Both specific conductivity and dissolved oxygen (DO) concentration decreased slightly at higher temperatures (Table 1; rmANOVA: specific conductivity, $F_{5,30} = 8.90$, $P < 0.001$; DO, $F_{5,30} = 86.24$, $P < 0.001$).

Biomass and community structure

Warming had a strong positive effect on biomass and chl *a* of stream biofilms. Biomass (as AFDM) varied 2.8-, 20-, and 24-fold across the temperature treatments at 30, 42, and 58 days, respectively (Fig. 1a, Table 2), while areal chl *a* concentrations varied 50-, 231-, and 174-fold at 30, 42 and 58 days respectively (Fig. 1b, Table 2). Biofilm community composition was dominated by N₂-fixing cyanobacteria (largely *Nostoc* and *Anabaena* spp.) in all temperature treatments (Table 3). This functional group comprised >91% of community biovolume regardless of temperature. The remaining functional groups included green and yellow-green algae (2.8 ± 8.0%, mean ± SD), diatoms (1.8 ± 1.5%), and diatoms with cyanobacterial endosymbionts (largely *Epithemia* and *Rhopalodia* spp.; 0.5 ± 1.4; Table 3).

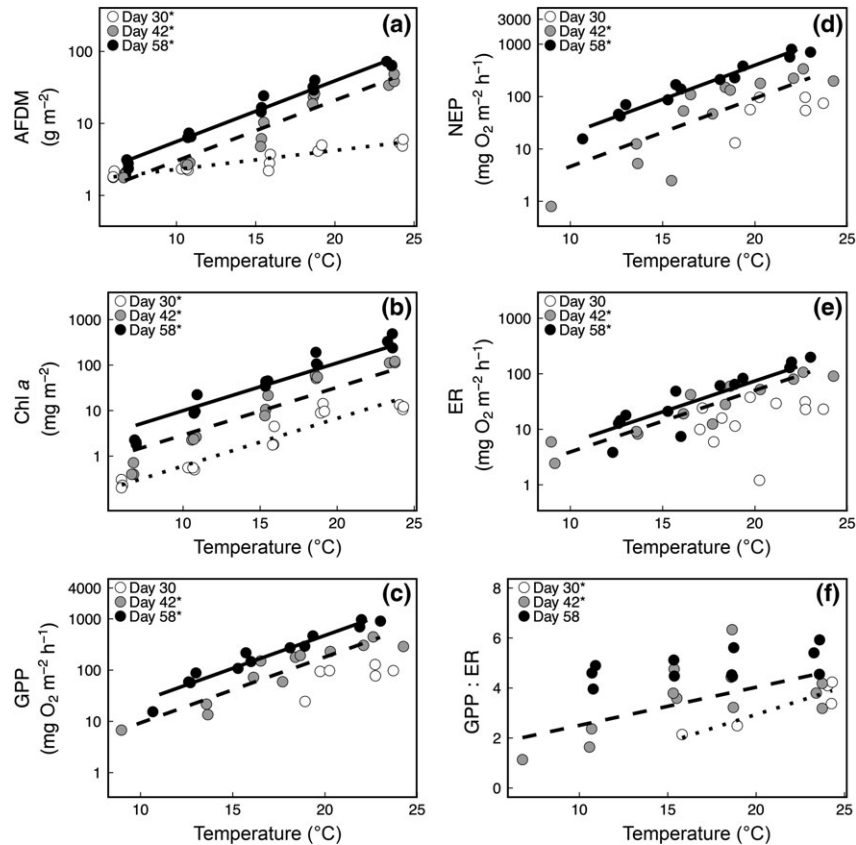


Fig. 1 Ash-free dry mass (AFDM); (a), chlorophyll *a* (chl *a*; b), gross primary production (GPP); (c), net ecosystem production (NEP); (d), ecosystem respiration (ER); (e), and the instantaneous (not diurnally integrated) ratio of GPP to ER (f) increased across the temperature gradient. Scale of *y*-axes is log_e-linear, and asterisks and model fit are displayed when significant ($\alpha \leq 0.05$). Dotted, dashed, and solid trend lines represent the first, second, and third sample periods respectively.

Metabolism

Rates of NEP, GPP, and ER exhibited a positive exponential relationship with temperature, excepting the first sampling period when limited data suggested similar trends, but precluded characterization of the relationship with temperature. On days 42 and 58, GPP varied 51- and 55-fold across the treatments, NEP varied 317- and 44-fold, and ER varied 22- and 42-fold respectively (Fig. 1c, d, e; Table 2). On day 30, GPP, NEP and ER did not vary between the two warmest treatments in which we were able to measure these fluxes. During daylight hours the channels were strongly autotrophic (GPP:ER = 4.0 ± 1.2 , mean \pm SD) and during the first two sampling periods, the degree of autotrophy increased with temperature (i.e. 2- and 5.5-fold over the temperature range; least-squares regression; 30 days; $r^2 = 0.80$, $P = 0.024$, intercept = -1.4 ± 1.1 , $\beta = 0.22 \pm 0.05$; 42 days; $r^2 = 0.29$, $P = 0.039$, intercept = 0.98 ± 1.1 , $\beta = 0.15 \pm 0.06$, estimate \pm SE; Fig. 1f).

Biofilm stoichiometry

Temperature had inconsistent effects on biofilm C and N content, except for a single sampling date (day 58) where significant increases with temperature were observed (Fig. 2a, b). Warming led to increased C (1.3-fold; from 13.9 to 25%) and N content (1.4-fold; from 1.6 to 2.8%) at 58 days (Fig. 2a, b; Table 4). On this date, C and N content was highest in the two middle temperature treatments (15.5 and 19 °C), relative to the coldest (7.5 and 11.2 °C) and warmest (23.6 °C) treatments (Table S1). Biofilm P content, which varied 2.0- to 2.5-fold across the treatments, exhibited a positive-saturating relationship with temperature (Fig. 2c; Table 4; Table S1). Biofilm C:N ratios did not differ significantly among temperature treatments, with the exception of a small, but significant decrease with temperature on a single date (30 days; Fig. 2d; Table 4; Table S1). In contrast, warming had a significant negative effect on biofilm C:P (30 and 58 days) and N:P ratios (30, 42, and 58 days). These changes were largely driven by relatively high C:P and N:P ratios in the coldest (i.e. ambi-

Table 2 Results of mixed-effects models for $\log_e(x)$ -transformed ash-free dry mass (AFDM, g m^{-2}), chlorophyll *a* (Chl *a*, mg m^{-2}), gross primary productivity (GPP, $\text{mg O}_2 \text{ m}^{-2} \text{ h}^{-1}$), net ecosystem productivity (NEP, $\text{mg O}_2 \text{ m}^{-2} \text{ h}^{-1}$), ecosystem respiration (ER, $\text{mg O}_2 \text{ m}^{-2} \text{ h}^{-1}$), ammonium ($\text{NH}_4^+\text{-N}$) uptake ($\mu\text{g N m}^{-2} \text{ h}^{-1}$), and soluble reactive phosphorus (SRP) uptake ($\mu\text{g P m}^{-2} \text{ h}^{-1}$).

Fixed effects	AFDM	Chl <i>a</i>	GPP	NEP	ER	$\text{NH}_4^+\text{-N}$ uptake	SRP uptake
	β (SE)	β (SE)	β (SE)	β (SE)	β (SE)	β (SE)	β (SE)
Intercept	0.22 (0.10)*	-2.91 (0.23)***	-2.01 (0.46)***	-2.61 (0.44)***	-2.45 (0.37)***	5.95 (0.23)***	5.51 (0.34)***
Temp.	0.06 (0.01)***	0.24 (0.01)***	0.29 (0.01)***	0.30 (0.01)***	0.25 (0.01)***	0.07 (0.001)***	0.07 (0.01)***
Day 42	-1.06 (0.13)***	1.56 (0.16)***	1.32 (0.22)***	1.14 (0.26)***	1.26 (0.35)**	0.67 (0.19)**	-0.39 (0.36)
Day 58	-0.45 (0.14)**	2.80 (0.17)***	2.29 (0.30)***	2.59 (0.31)***	1.64 (0.34)***	0.61 (0.18)**	-0.09 (0.29)
Temp.*	0.13 (0.01)***	–	–	–	–	–	–
Day 42							
Temp.*	0.13 (0.01)***	–	–	–	–	–	–
Day 58							
Random effects	SD	SD	SD	SD	SD	SD	SD
Channel (intercept)	0.14	0.43	0.70	0.79	0.97	0.63	0.94
Day 42 (slope)	0.16	0.60	0.53	0.30	1.22	0.67	1.28
Day 58 (slope)	0.19	0.64	0.87	0.87	1.17	0.56	0.94
Residuals	0.05	0.17	0.14	0.19	0.15	0.16	0.14
$R^2_{\text{GLMM(m)}}$	0.96	0.91	0.85	0.73	0.68	0.57	0.29
$R^2_{\text{GLMM(c)}}$	0.99	0.99	0.98	0.98	0.98	0.94	0.97

We initially fit a 'full' model that included temperature, day, and a temperature \times day interaction as fixed effects and channel identity as an additional random effect. We used model selection techniques to select the most appropriate model (Akaike Information Criterion; AIC). The marginal R^2 (GLMM[m]; fixed effects only) and the conditional R^2 (GLMM[c]; fixed and random effects) represent the proportion variance explained by each model. Day is approximate, as response variables were measured on separate, but generally consecutive days. SE = standard error; SD = standard deviation; * $P < 0.05$, ** $P < 0.01$, *** $P < 0.001$.

Table 3 Functional group composition of biofilm communities (mean % \pm 1 SD) by biovolume and averaged across time periods

Mean temperature	Cyanobacteria	Diatoms	N_2 -fixing diatoms	Green and yellow-green algae
7.5	98.2 (2.3)	1.6 (2.3)	0.02 (0.07)	0.04 (0.08)
11.2	97.9 (1.7)	1.1 (0.4)	0	0.8 (1.6)
15.5	98.5 (1.0)	1.2 (1.0)	0.02 (0.03)	0.2 (0.3)
19.0	91.7 (14.8)	2.4 (1.7)	0.7 (1.6)	5.1 (13.3)
23.6	91.6 (8.5)	2.9 (1.6)	1.3 (2.0)	4.0 (7.2)

ent) temperature treatment (Fig. 2e, f; Table 4; Table S1). Warming above ambient thermal conditions produced a 1.9-fold (C:P) and 1.8-fold (N:P) decrease in elemental ratios; however, no differences were found among treatments above ambient temperature (Fig. 2e, f; Table 4; Table S1).

Nutrient uptake and nitrogen fixation

Rates of $\text{NH}_4^+\text{-N}$ uptake and N_2 -fixation increased across the temperature gradient (Fig. 3a, b; Tables 2 and 4) and, in the warmest treatments, N_2 -fixation became the dominant flux of N to the biofilm (Fig. 3c).

Temperature had a strong positive effect on N_2 -fixation rates, which varied 22-, 121-, and 64-fold on days 41, 53, and 69 (Fig. 3b; Table 4). $\text{NH}_4^+\text{-N}$ uptake rates were positively and exponentially related to temperature, but the increase across the thermal gradient was much shallower than that of N_2 -fixation, increasing only 6.8-, 3.2-, and 2.8-fold on days 31, 44, and 65 (Fig. 3a; Table 2). Thus, in the cold treatments, $\text{NH}_4^+\text{-N}$ uptake rates were greater than N_2 -fixation rates, especially early in the experiment. In the warmest treatments, N_2 -fixation rates were over two times higher than $\text{NH}_4^+\text{-N}$ uptake rates (Fig. 3c). SRP uptake rates varied 8.6-fold across all treatments and dates, but were only

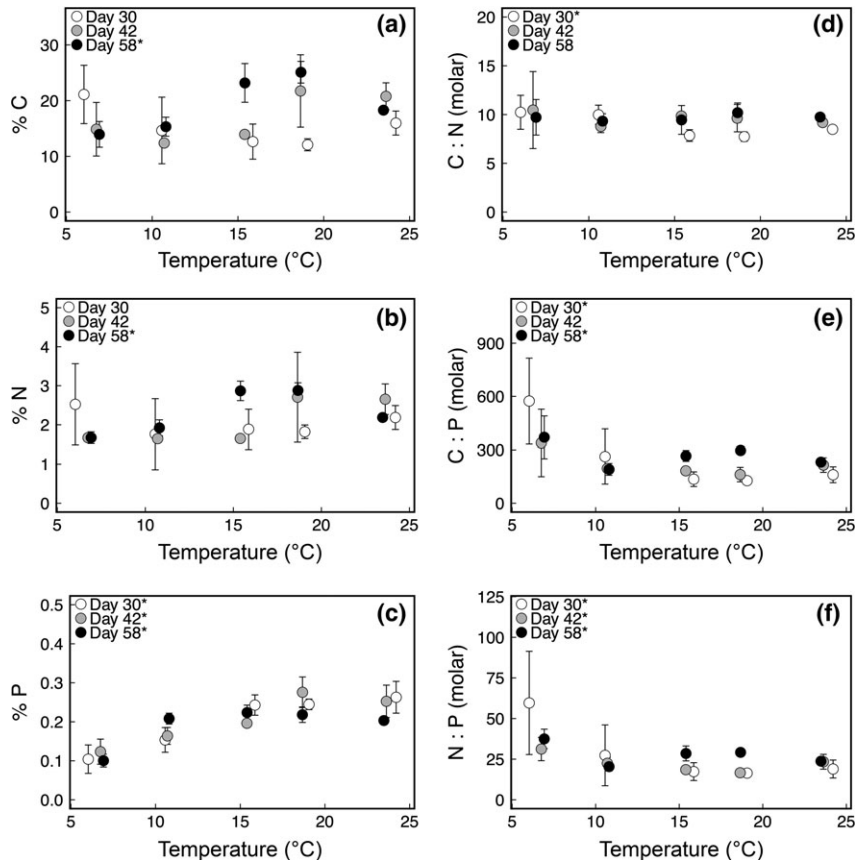


Fig. 2 Carbon (C, a) and nitrogen (N, b) content showed mixed nonlinear responses to temperature that were generally positive, while phosphorus (P, c) content clearly increased with temperature. C:N (d) was temperature-invariant, whereas C:P (e) and N:P (f) ratios decreased at temperatures above ambient. Asterisks displayed in the legend indicate significant relationships ($\alpha = 0.05$), error bars are ± 1 SD.

Table 4 Repeated measures analysis of variance was used in instances where residual plots suggested evidence of a nonlinear temperature response

	Log _e N ₂ -fixation ($\mu\text{g N m}^{-2} \text{h}^{-1}$)	C (%)	N (%)	P (%)	C:N (molar ratio)	C:P (molar ratio)	N:P (molar ratio)
Intercept	—	—	—	—	—	—	—
Degrees of freedom	1, 28	1, 20	1, 20	1, 28	1, 20	1, 20	1, 20
F-statistic	14513.7***	2857.9***	2698.5***	2718.6***	7096.3***	2203.6***	1957.4***
Treatment (temp.)	—	—	—	—	—	—	—
Degrees of freedom	4, 10	4, 10	4, 10	4, 10	4, 10	4, 10	4, 10
F-statistic	145.6***	27.6***	24.15***	40.5***	4.1*	15.6***	9.9**
Day	—	—	—	—	—	—	—
Degrees of freedom	2, 28	2, 20	2, 20	2, 28	2, 20	2, 20	2, 20
F-statistic	11.8***	6.7**	2.4	0.8	2.0	1.2	9.3**
Treatment (temp.) * Day	—	—	—	—	—	—	—
Degrees of freedom	—	8, 20	8, 20	—	8, 20	8, 20	8, 20
F-statistic	—	6.6***	3.7**	—	1.2	2.9*	3.1*

A full model was fit for each variable with the covariates temperature, day, and a temperature \times day interaction. Model selection methods (Akaike Information Criteria; AIC) were used to determine the final model. Carbon (C) content, nitrogen (N) content, C:N, C:phosphorus (P), and N:P were all fit with the full model, and N₂-fixation and P content were fit with a reduced model containing only the terms temperature and day. * $P < 0.05$, ** $P < 0.01$, *** $P < 0.001$.

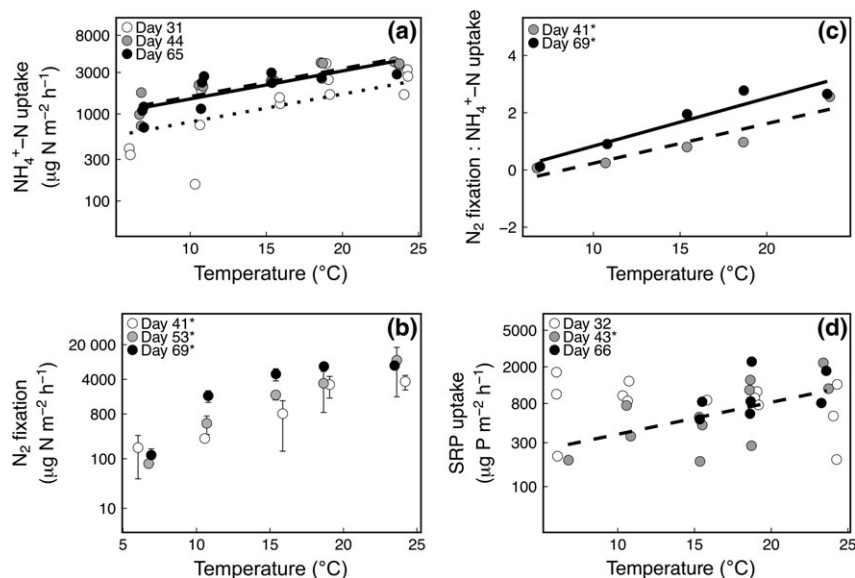


Fig. 3 Rates of nitrogen uptake ($\text{NH}_4^+\text{-N}$); (a) and N_2 -fixation (b) increased across the temperature gradient. At warmer temperatures, N_2 -fixation supplied the majority of total N (c) flux to the system. Phosphorus (P) uptake (as soluble reactive phosphorus; SRP) increased with temperature on a single date and was highly variable (d). Model fit and asterisks displayed when significant ($\alpha = 0.05$). Scale of y -axes in panels a,b,d are \log_e -linear, and dotted, dashed, and solid lines represent the first, second, and third sample periods respectively. Ratios of N_2 -fixation:N uptake utilize data measured on separate dates under similar conditions. Error bars are ± 1 SD.

weakly related to temperature on one sampling date (Fig. 3d; Table 2).

Predicted nutrient demand

We compared our measured N (i.e. $\text{NH}_4^+\text{-N}$ uptake and N_2 -fixation) and P (i.e. SRP uptake) assimilation rates to predicted nutrient demand based on GPP and biofilm stoichiometry (Fig. 4). SRP uptake was positively related ($r^2 = 0.34$, $P = 0.003$) to predicted P demand, and the 95% confidence interval for this relationship overlapped the 1:1 line for most of the range in predicted P demand; however, this was not true at the lower end of the predicted values where the majority of the data points were located (Fig. 4a). $\text{NH}_4^+\text{-N}$ uptake was similarly related to predicted $\text{NH}_4^+\text{-N}$ demand ($r^2 = 0.36$, $P = 0.032$; Fig. 4b), yet results from treatments warmer than ambient temperatures fell far below the 1:1 line, suggesting a decoupling between GPP and $\text{NH}_4^+\text{-N}$ uptake with warming. Including N_2 -fixation in our measurement of total N assimilation (on the y -axis) resulted in a much greater correspondence between predicted and measured N assimilation (Fig. 4c). Nonetheless, we still tended to over-predict N assimilation at cold temperatures and under-predict N assimilation at warm temperatures (Fig. 4c).

Nutrient use efficiency

We examined nutrient use efficiency (UE) from two perspectives: (i) bulk stoichiometry, which represents

net accumulation of elements, and (ii) relative fluxes of C and nutrients into biofilms. C:P stoichiometry varied little across the thermal gradient, suggesting that P use efficiency (P-UE) was unaffected by temperature. However, estimates of P-UE based on the fluxes of C and P to the biofilm suggested that warming had a positive effect on P-UE on day 42, the single date containing sufficient data (a 37-fold increase; $r^2 = 0.70$, $P = 0.046$, intercept = 0.08 ± 1.15 , $\beta = 0.23 \pm 0.07$, $\log_e \pm \text{SE}$; Fig. 5a). The generally invariant response of C:N stoichiometry across temperature treatments suggests that N use efficiency (N-UE) was not influenced by temperature using this metric. In contrast, N-UE estimates based on relative fluxes of C and N to biofilms indicated that temperature had a generally positive effect on N-UE. We found that N-UE calculated with $\text{NH}_4^+\text{-N}$ uptake rates only (N-UE_U) increased 56-fold across the temperature gradient (42 days; $r^2 = 0.87$, $P = 0.012$, intercept = -2.81 ± 0.85 , $\beta = 0.28 \pm 0.05$; 58 days; $r^2 = 0.98$, $P < 0.001$, intercept = 0.47 ± 0.14 , $\beta = 0.16 \pm 0.009$, $\log_e \pm \text{SE}$; Fig. 5b). N-UE_{UF} (i.e. including both N_2 -fixation and $\text{NH}_4^+\text{-N}$ uptake as N fluxes to the biofilm) also indicated that warming led to increased nutrient UE, but the slopes of these relationships were shallower (average of a 17-fold increase; 42 days: $r^2 = 0.75$, $P = 0.035$, intercept = -2.36 ± 0.94 , $\beta = 0.21 \pm 0.05$; 58 days: $r^2 = 0.72$, $P = 0.042$, intercept = 0.68 ± 0.43 , $\beta = 0.09 \pm 0.02$, $\log_e \pm \text{SE}$; Fig. 5c).

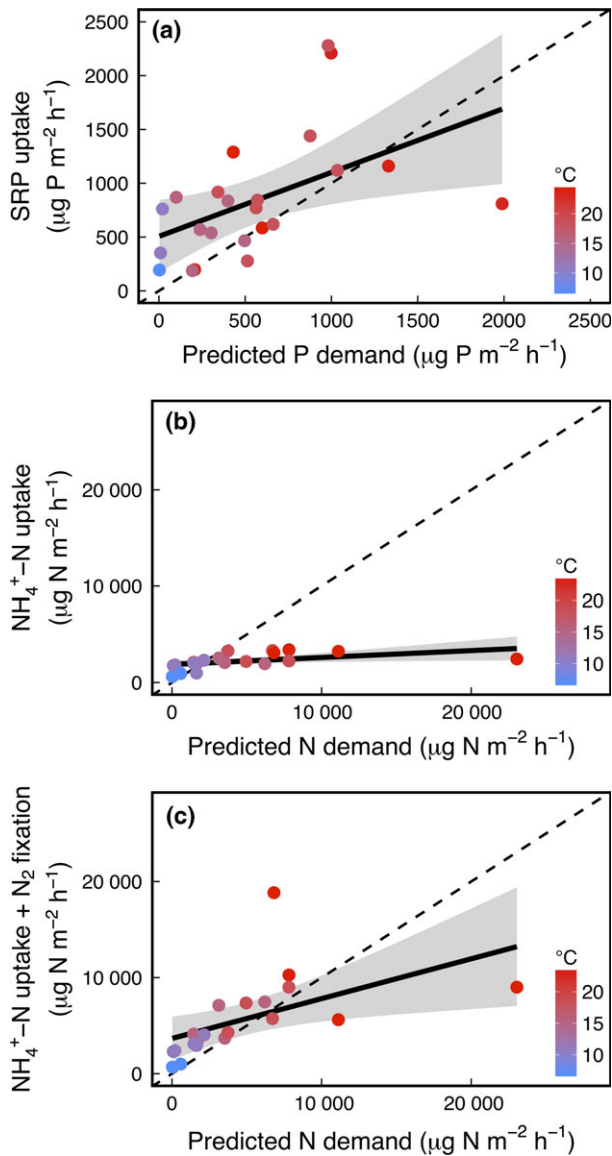


Fig. 4 Phosphorus (P, a) uptake generally exceeded predicted P demand, whereas NH_4^+ -N uptake (b) only exceeded N demand at relatively cold temperatures. Incorporating N_2 -fixation in estimates of total N acquisition (NH_4^+ -N uptake + N_2 -fixation); (c) led to a closer match between total measured and predicted N demand. Dashed line is 1:1 slope, solid line is least squares regression with 95% confidence intervals in gray.

Discussion

Our controlled warming experiment revealed strong effects of temperature on the ecosystem function of stream biofilms. Metabolic processes, including gross primary production, ecosystem respiration, and net ecosystem production all increased with warming, despite the relatively low supply of ionic nitrogen known to be limiting at our study site (Friebert *et al.*,

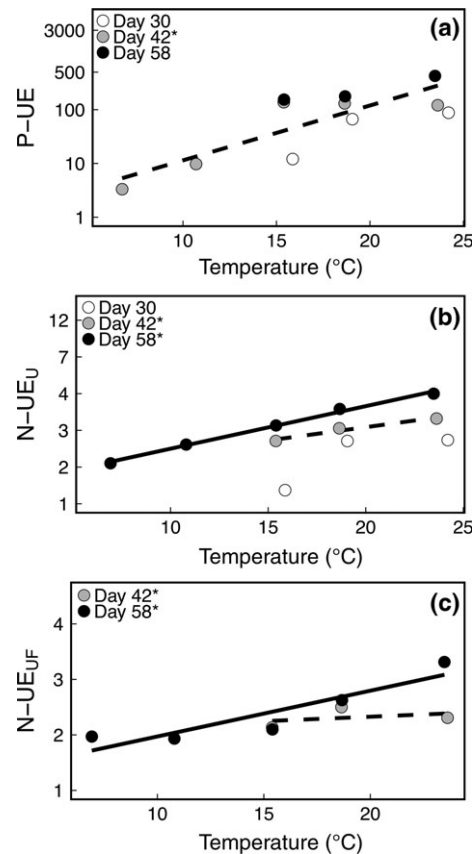


Fig. 5 Phosphorus (P-UE, a), NH_4^+ -N (N-UE_U, b), and N_2 -fixation + NH_4^+ -N (N-UE_{UF}, c) use efficiencies increased with increasing temperature when calculated using the relative fluxes of carbon and nutrients (see text). Asterisks and least-squares regressions indicate significant relationships ($\alpha = 0.05$). Scales of y-axes are log_e-linear, and dashed and solid trend lines represent the second and third sample periods respectively. Sampling dates in the legend are approximate because the fluxes used to calculate nutrient use efficiency (NEP, uptake, and N_2 -fixation) were measured on separate days.

2009). The biomass of autotrophs also increased substantially across the thermal gradient, contrary to recent experiments conducted in lentic environments (Yvon-Durocher *et al.*, 2010b; Dossena *et al.*, 2012; Shurin *et al.*, 2012; *but see* Baulch *et al.*, 2005), as well as predictions based on metabolic theory (Allen *et al.*, 2005; Cross *et al.*, 2015). Unexpectedly, biofilm C:nutrient ratios showed a null or negative response (i.e. no change in C:N and reduced C:P ratios) to warming, suggesting important temperature-induced changes in how nutrients are acquired, stored, and routed. Our results shed new light on the complex ways in which environmental warming may alter the coupling of carbon and nutrient cycles in stream ecosystems.

We found that biofilm C:N ratios remained relatively constant across the temperature treatments,

despite a significant increase in ecosystem metabolism. These results suggest that both N and P assimilation 'kept pace' with C fixation rates as biomass accumulated with warming. With respect to N, this result was especially surprising given its very low concentration in the source stream. Although ammonium uptake increased significantly with temperature, this flux was grossly insufficient to meet the total predicted N demand (i.e. based on GPP and biomass stoichiometry; Webster *et al.*, 2003). There are a few possible explanations for this pattern. First, it is possible that increased N demand at warmer temperatures was met by elevated rates of N₂-fixation, bringing a new source of N to biofilms in the form of N₂ gas. Our measured rates of N₂-fixation strongly support this hypothesis and suggest that warming can alter the dominant sources and pathways of nitrogen in stream ecosystems (Welter *et al.*, 2015). In the coldest treatment N₂-fixation supplied ~9% of total N assimilation and ~72% in the warmest treatment, effectively shifting the primary N source from dissolved ammonium uptake to N₂-fixation as the primary N source. If such a pattern is general (Paerl & Huisman, 2008), more effort should be dedicated to understanding how warming-induced shifts in the N cycle may alter the balance of N uptake and removal (Mulholland *et al.*, 2000), as well as the flux of N through higher trophic levels (Marcarelli *et al.*, 2008). Moreover, our results highlight the importance of considering functional traits (e.g., N₂-fixation) when predicting how coupled biogeochemical cycles will respond to climate change (Webb *et al.*, 2010; Finzi *et al.*, 2011).

Secondly, uptake of nitrate-N may be important for supporting biofilm growth, and we did not explicitly measure this flux. Measurements of whole-stream ammonium vs. nitrate uptake in 9 Hengill streams that vary across a similar temperature range show that ammonium uptake is generally higher than that of nitrate (ratios from ~2:1 to 50:1), consistent with other recent studies in these streams (J.M. Hood, unpublished data; Demars *et al.*, 2011; Rasmussen *et al.*, 2011). Despite this, it seems likely that under conditions of higher demand (i.e. warmer temperatures) nitrate-N assimilation would become increasingly important, especially considering the relatively higher energetic cost of N₂-fixation (Howarth *et al.*, 1988). This unaccounted source of N may explain why we progressively over-predict N assimilation at warmer temperatures.

We also found that biofilms became relatively more P-rich (reduced C:P ratios) in response to warming, in contrast with our original predictions. Although unexpected, this result is not inconsistent with recent work by De Senerpont Domis *et al.* (2014), which

showed little response of algal C:P stoichiometry with warming under P-rich conditions. In our study area, dissolved P is relatively abundant (i.e. low dissolved N:P; O'Gorman *et al.*, 2012) and this may help explain the lack of a strong SRP uptake response to warming. Additionally, our channels lacked macrophytes which have been shown to strongly influence the relationship between temperature and P uptake in natural streams (Demars *et al.*, 2011). Indeed, our estimates of predicted P demand from GPP and biofilm stoichiometry suggest that uptake of dissolved P roughly met, or exceeded, expected P demand at all experimental temperatures.

The microbial assemblage that developed on tiles was dominated by globally ubiquitous cyanobacteria (Paerl & Huisman, 2008; Lürling *et al.*, 2013). A shift from *Nostoc* spp. toward *Anabaena* and *Cylindrospermum* spp. was observed across the thermal gradient (J.R. Welter & P.C. Furey, St. Catherine University, unpublished data). Such a shift in assemblage structure is likely due to thermal optima of individual cyanobacteria taxa, as *Anabaena* exhibits a relatively high temperature optimum (Kratz & Myers, 1955; Lürling *et al.*, 2013) and was only observed in the warmest treatment. Changes in species composition were accompanied by increases in N₂-fixation rates with warming, which may represent a direct response of the nitrogenase enzyme to increasing temperature (Ceuterick *et al.*, 1978), as well as interspecific differences in N₂-fixation potential among taxa. N₂-fixers typically have high P requirements, which may result from the energy-intensive, high adenosine triphosphate (ATP) cost of N₂-fixation (Vitousek *et al.*, 2002). Such high P requirements, together with changes in species composition, may explain the slight reduction in C:P stoichiometry with warming.

Although our prediction that warming would lead to elevated C:nutrient ratios was not supported, we did find a strong positive effect of warming on flux-based nutrient UEs. Such a result could arise from increased nutrient UE of individuals (i.e. physiological changes; Baligar *et al.*, 2001) or community-level shifts toward taxa that use nutrients more efficiently at warm temperatures (Hiremath & Ewel, 2001), neither of which would necessarily be reflected in the elemental composition of biofilms. Indeed, the metabolic demand for nutrients may not be clearly reflected in biomass if some elements, such as N, are bound to recalcitrant organic molecules and not metabolically active. Moreover, large variation exists in cellular activity within biofilms (Lock *et al.*, 1984) and, if the majority of nutrient uptake takes place on the surface among actively growing cells (Burkholder *et al.*, 1990), we might expect flux-based nutrient UE to increase as bio-

films develop and thicken. This may result from internal recycling processes, rather than water column uptake, supplying an increasing proportion of nutrient demand. Finally, part of the discrepancy between patterns of biomass stoichiometry and nutrient UE may be methodological, stemming from a potential disconnect between using pools (i.e. stoichiometry) vs. fluxes (i.e. rates) to assess nutrient UEs. Our estimates of nutrient UE were derived from instantaneous measurements of NEP and nutrient uptake during daylight hours when both fluxes are predicted to be high. Yet biomass stoichiometry measurements integrate element cycling over much longer time scales, suggesting that measures of nutrient UE may be best integrated over time periods that match the scale at which elements change in biomass (e.g., litterfall vs. leaf C:nutrient ratios; Vitousek, 1982). Clearly, additional research is needed to disentangle relationships between biomass stoichiometry and nutrient UE, as well as expectations about how warming and diel patterns of nutrient acquisition might influence nutrient UEs.

Anticipating how climate change affects stream ecosystems will require a broad array of research techniques that draw from theory, observation, and experimentation. Results from our experiment demonstrate that temperature had a strong influence on benthic biofilm structure and function, as well as coupled elemental cycles; however, it is unclear how these results apply to broader spatial and temporal scales (e.g. years to centuries; Osmond *et al.*, 2004). Recent complementary research across a natural gradient of stream temperatures at Hengill suggests that our results generally parallel warming responses at larger, whole-stream scales. For instance, increased metabolism with warming has been observed along the natural thermal gradient (Demars *et al.*, 2011; Perkins *et al.*, 2012) and natural biofilm stoichiometry also appears to be relatively temperature-invariant (J.R. Junker & W.F. Cross, Montana State University, unpublished data). These patterns suggest that results from our small-scale, short-term warming experiment may indeed predict large-scale, long-term responses to warming at this location. Nonetheless, additional studies are needed that explicitly compare the responses of short-term and small-scale experiments to observations across natural gradients of temperature or time (Dunne *et al.*, 2004; O'Gorman *et al.*, 2012). Geothermally heated ecosystems may be particularly well-suited for such investigations (O'Gorman *et al.*, 2014).

We have demonstrated that warming can strongly influence the structure and function of biofilms in aquatic ecosystems. Our findings suggest that rates of C-based fluxes in these habitats will increase with warm-

ing due to a combination of decreased inhibition by cold temperature and increased N availability. However, our results also suggest that how these responses translate to other elemental cycles (e.g., N and P) will depend upon environmental context (e.g., nutrient supply) and the functional traits of existing taxa. In oligotrophic streams, such as those studied here, the traits of dominant functional groups (i.e. cyanobacteria) may interact with ambient nutrient availability to strongly influence how the coupling of C to other key elements will respond to warming. Future studies should simultaneously consider species functional traits, thermal performance, and temperature dependent variation in resource availability when studying effects of warming on fluxes of energy and elements in aquatic ecosystems.

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Supporting Information

Additional Supporting Information may be found in the online version of this article:

Table S1. Results from a one-way ANOVA with Tukey's honestly significant difference (HSD) test of pairwise temperature treatment comparisons.

Figure S1. Time series of daily mean temperature treatments.

Figure S2. Experimental channel array in operation.